

# A NEW EOCENE CASSIGERINELLA FROM FLORIDA

by W. G. CORDEY

ABSTRACT. A new species of *Cassigerinella* (*C. eocaenica*) is proposed for specimens obtained from an Eocene sample taken off Blake Plateau, Florida. The species is similar, but quite distinct from the Oligocene marker *C. chipolensis* (Cushman and Ponton 1932). Therefore the stratigraphical value of the overlap in ranges of *C. chipolensis* with *Pseudohastigerina micra* (Cole 1927) as indicators of basal Oligocene is not affected.

THE joint occurrence of the species *Pseudohastigerina micra* (Cole 1927) and *Cassigerinella chipolensis* (Cushman and Ponton) has been regarded as a means of recognizing the basal Oligocene, at least within the tropical and semi-tropical belts. This overlap was first recognized by Blow and Banner (*in* Eames *et al.* 1962, p. 68) as a further criterion for the recognition of their zone of *Globigerina sellii* Borsetti 1959 (= *G. oligocaenica* Blow and Banner 1962). Saunders and Cordey (1965) also recognized this overlap in a study of the Oceanic Formation of Barbados. It occurred in samples immediately overlying deposits of definite Eocene age containing such forms as *Hantkenina* spp., *Globorotalia centralis*, and *G. cerro-azulensis*. Bolli (1966), in a general review of world-wide planktonic zonation, added further support to the stratigraphic value of this observed overlap.

Saito and Bé (1963) established an Oligocene age (*sensu* Eames *et al.*, *op. cit.*) for the Vicksburg group of the Gulf Coast region. However, they stated (p. 704) that *C. chipolensis* and *P. micra* occurred together with Eocene forms (e.g. *Hantkenina alabamensis*, *H. primitiva*, *Globorotalia cerro-azulensis*) from a core taken off the Florida coast (Lamont Core A167-21, 29° 49' N., 79° 39' W.). This record therefore cast considerable doubt on the stratigraphic value of this overlap as far as the recognition of the Eocene-Oligocene boundary or basal Oligocene was concerned. Blow examined these cassigerinellids and concluded (*pers. comm.*) that they were different from *C. chipolensis*. Through the kindness of Drs. Bé and Saito the writer obtained material which contained the same form of *Cassigerinella* with *Hantkenina* spp., but from a location to the north of the Lamont Core, at Blake Plateau, 30° 04.8' N., 79° 14.5' W., Sample J.6B.

A careful comparison of the cassigerinellids in the Blake Plateau material with *C. chipolensis* from both the *Globigerina ampliapertura* zone of Trinidad and the Oceanic Formation of Barbados indicates that the Eocene specimens differ from *C. chipolensis*. In view of the stratigraphic significance of *C. chipolensis*, it is desirable that a new species should be erected for these Eocene occurrences.

## SYSTEMATIC DESCRIPTIONS

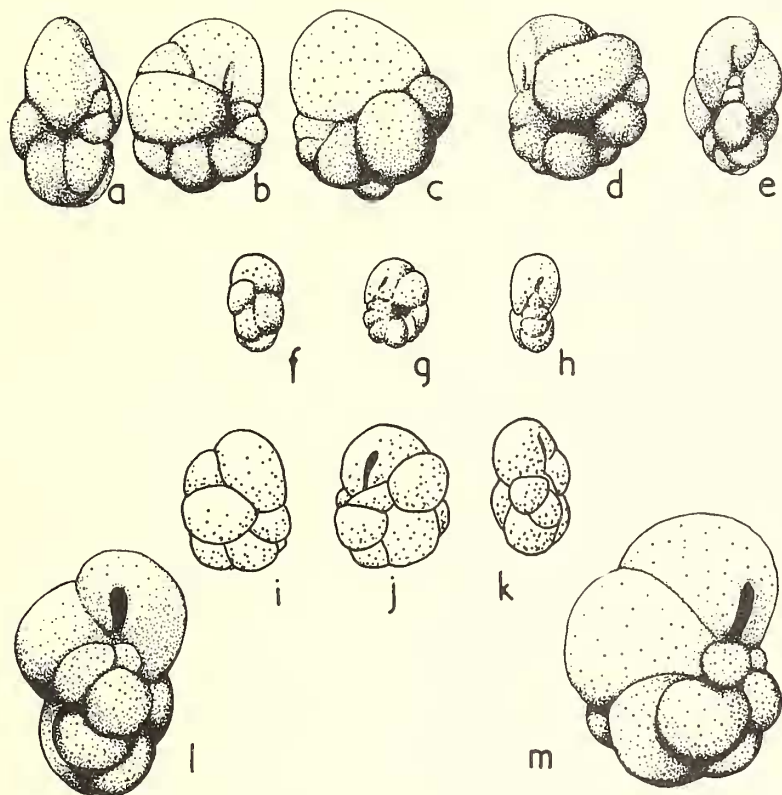
Genus CASSIGERINELLA Pokorný 1955

*Cassigerinella eocaenica* sp. nov.

Text-fig. 1a-e

*Description.* Test very small, calcareous, perforate; 8-9 chambers in the final whorl, showing a gradual and uniform increase in size, moderately inflated. Initial two (possibly

three) chambers planispirally arranged, later becoming biserial alternating, but coiled in the same plane. Periphery lobulate, initially sub-acute, becoming sub-rounded, sutures distinct, depressed, curved. Aperture a latero-marginal, extra-umbilical arch.



TEXT-FIG. 1.

*a-e*, *Cassigerinella eocaenica* sp. nov. Sample J-6b, 30° 04·8' N., 79° 14·5' W., 215' from top. *a-c*, Holotype, BMNH P46838; *d, e*, Paratype, BMNH P46839; both  $\times 300$ .

*f-m*, *Cassigerinella chipolensis* (Cushman and Ponton). *f-h*, Dissected hypotype showing penultimate whorl, early chambers planispirally coiled, later chambers alternating; *G. ampliapertura* Zone, Ciperó Coast, Trinidad; BMNH P46840,  $\times 213$ . *i-k*, After Cushman and Ponton,  $\times 150$ . *l, m*, Hypotype, *G. ampliapertura* zone, Ciperó Coast, Trinidad; BMNH P46841,  $\times 275$ .

*Remarks.* *C. eocaenica* is similar to *C. chipolensis*, but differs in being consistently smaller, the greatest breadth of the final whorl varying from 0·1 to 0·12 mm., the average being 0·1 mm. Measurement of 45 specimens of *C. chipolensis* shows a variation in the breadth of the final whorl from 0·13 to 0·19 mm., the average being 0·16 mm. Secondly, the chambers in the final whorl are less inflated than in *C. chipolensis*. Specimens of *C. chipolensis* from Barbados (Oceanic Formation), from the Ciperó Formation of Trinidad and from the Oligocene of Lindi (Blow and Banner in Eames *et al.* 1962, pl. 15, figs. M-N), show a more rapid increase in the size of the chambers of the final whorl

than in *C. eocaenica*. *C. chipolensis* also has a more rounded periphery throughout. Finally, the chamber arrangement in the final whorl of *C. eocaenica* varies from an initial planispiral arrangement to alternating, whereas the chamber arrangement in *C. chipolensis* is entirely alternating throughout the final whorl. The planispiral arrangement is only developed in the first two or three chambers of the penultimate whorl of *C. chipolensis* (text-fig. 1, *f-h*). The only other species which shows any morphological similarity to *C. eocaenica* is *C. globolocula* Ivanova 1958. Pokorný (*in Eames et al.*) agreed that his species *C. boudecensis* was probably conspecific with *chipolensis*, and considered that *globolocula* was 'certainly conspecific with *boudecensis*' (*op. cit.*, p. 83). The writer would agree with this conclusion and therefore the above remarks on *C. chipolensis* and *C. eocaenica* apply equally to *C. globolocula* (and also *C. boudecensis*).

*Deposition of types.* Holotype and paratype specimens are deposited in the British Museum (Natural History). An unfigured paratype is deposited in the United States National Museum, No. 643514.

*Material.* Fifteen specimens of *C. eocaenica*; forty-five specimens of *C. chipolensis*.

*Acknowledgements.* The author is grateful to Drs. Bé and Saito (Lamont Geological Observatory, Columbia University, New York) for the donation of the samples upon which this study is based; Dr. R. Lagaaij and J. A. Postuma (Bataafse Internationale Petroleum Maatschappij N.V., The Hague), for their critical reading of the manuscript; and the Bataafse Internationale Petroleum Maatschappij N.V., for permission to publish this paper.

#### REFERENCES

- BOLLI, H. M. 1966. Zonation of Cretaceous to Pliocene marine sediments based on planktonic Foraminifera. *Boln inf. Asoc. Venezolana de Geología, Mineraria y Petróleo*, 9(1), 3-32.
- EAMES, F. E. *et al.* 1962. *Fundamentals of Mid-Tertiary Stratigraphical Correlation*. Cambridge University Press.
- IVANOVA, L. G. *in* BYKOVA, N. K. 1958. Novye Rody i Vidy Foraminifera. *Trudy VNIIGRI*, no. 115, Mikrofauna SSR, 9, 4-81.
- SAITO, T. and BÉ, A. W. H. 1963. Planktonic Foraminifera from the American Oligocene. *Science*, 145, 703-4.
- SAUNDERS, J. B. and CORDEY, W. G. 1965. The biostratigraphy of the Oceanic Formation in the Bath Cliff section, Barbados. *Proc. 4th Caribbean Geol. Congress, Port of Spain, 1965*. (In press.)

W. G. CORDEY  
Bataafse Int. Petr. Mij.  
EP/12, Carel van Bylandtlaan 30  
The Hague  
Netherlands

Typescript received from author 29 June 1967